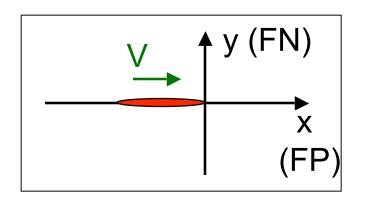
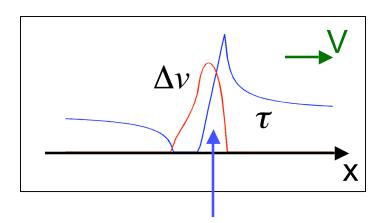


## Ground Motion from Dynamic Rupture Models

Steady-state slip pulse in 2D

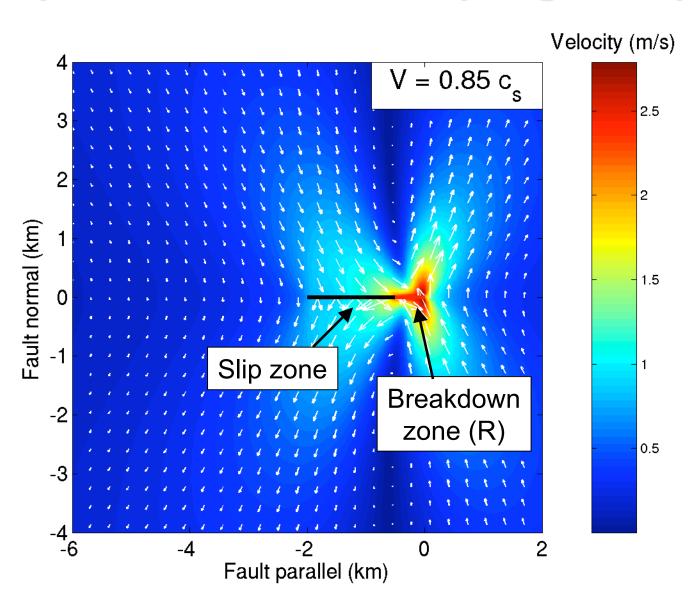




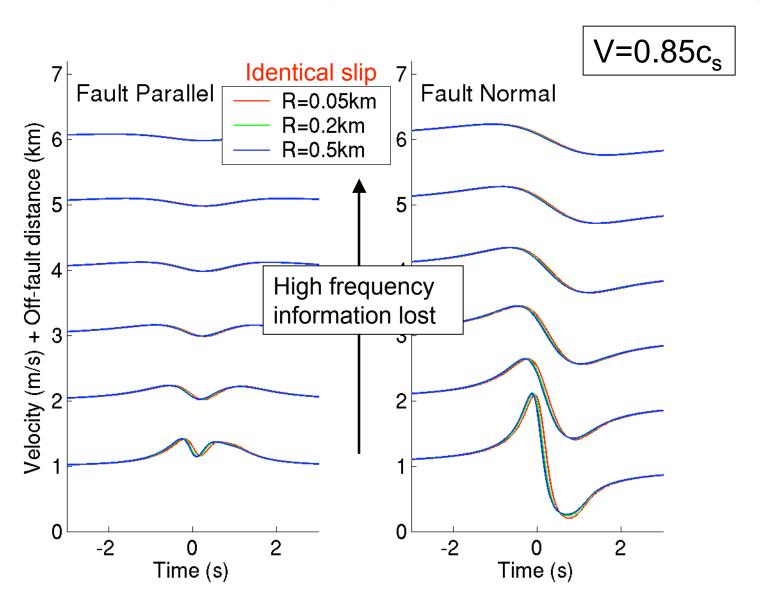
Dynamic (specify shear traction in slip zone)

Analytical expressions for velocity and stress fields, both on and off of the fault, have been derived (Broberg, 1978; Freund, 1979; Broberg, 1989; Rice et al., 2004; Dunham, 2004)

#### Velocity Field for Sub-Rayleigh Rupture



### Synthetic Seismograms: Sub-Rayleigh



### What properties of the rupture process are measurable?

More than a few km from the fault, ground motion is only sensitive to:

- 1. rupture speed V
- Kinematic 2. slip zone length L (or rise time) 3. final slip

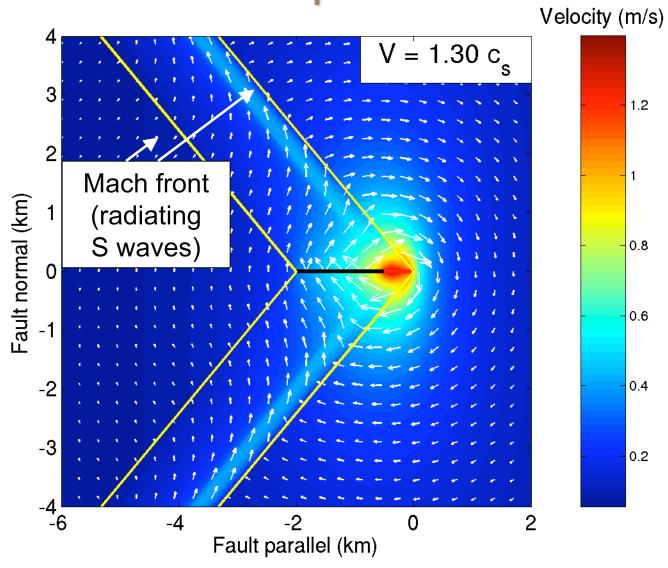
  - 4. fracture energy G (indirectly...)

#### Not sensitive to:

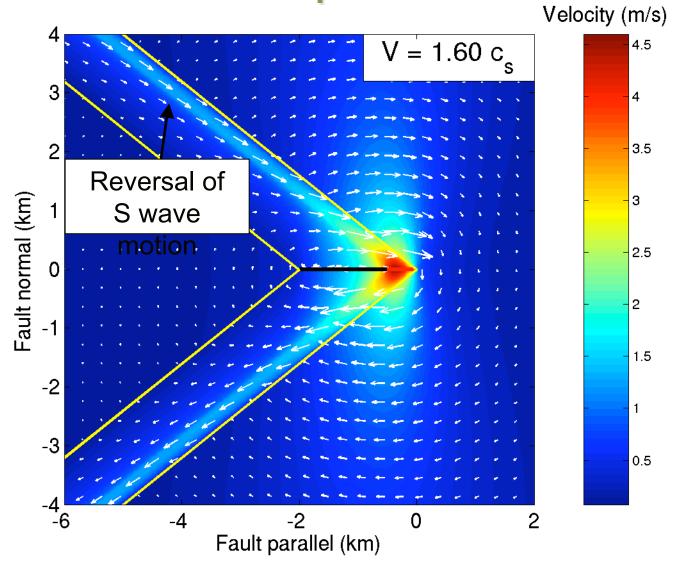
1. breakdown zone length R

This explains why kinematic models have been so successful and is bad news for seismologists interested in dynamics...

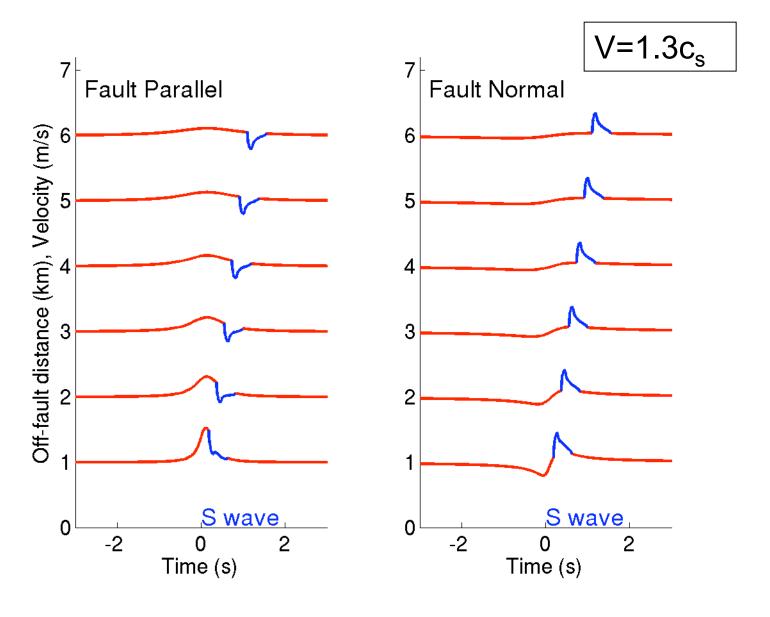
# Velocity Field for Intersonic Rupture



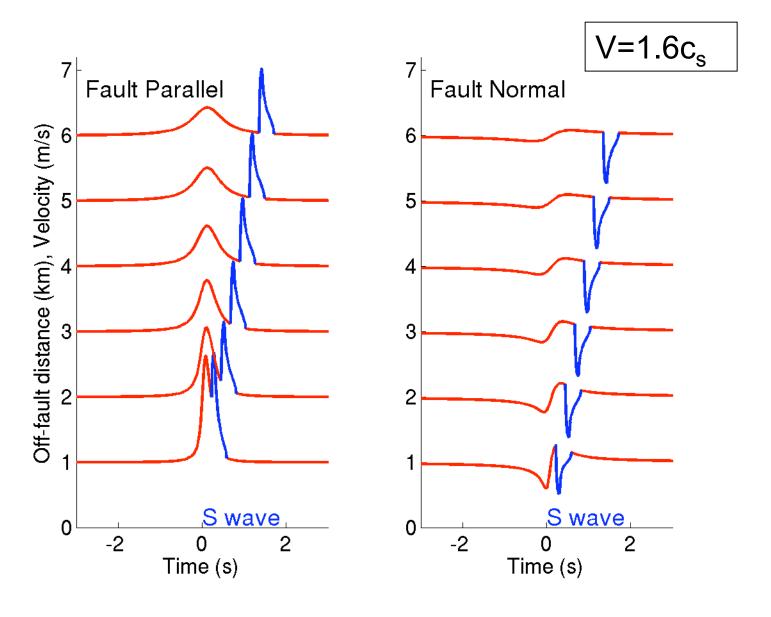
## Velocity Field for Intersonic Rupture



## Observational Signature



## Observational Signature



#### One can analytically show that:

Every S wave velocity and stress component recorded away from the fault traces the exact slip velocity history on the fault!

→ Very different attenuation relationships

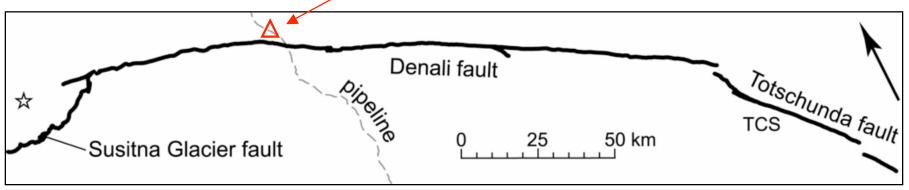
for supershear ruptures (Dunham, submitted to *GRL*, 2004) open-source Fortran code to be released soon

## 2002 Denali Fault Earthquake

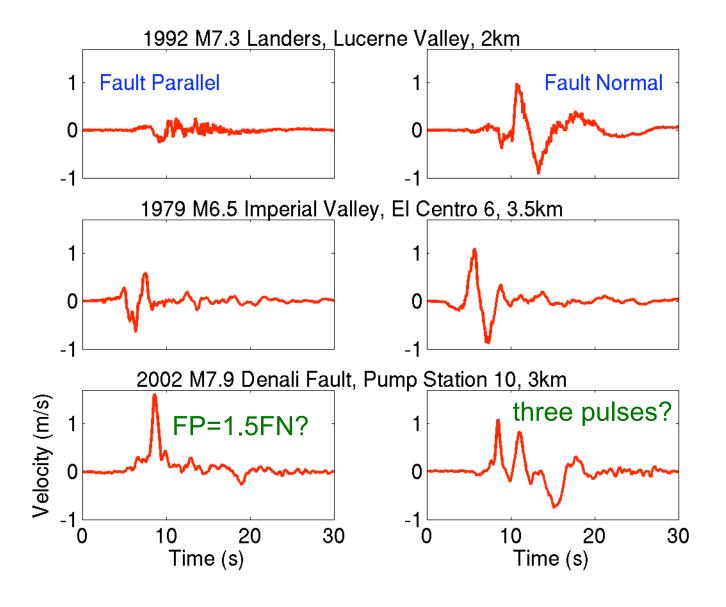


Pump Station 10 (PS10)

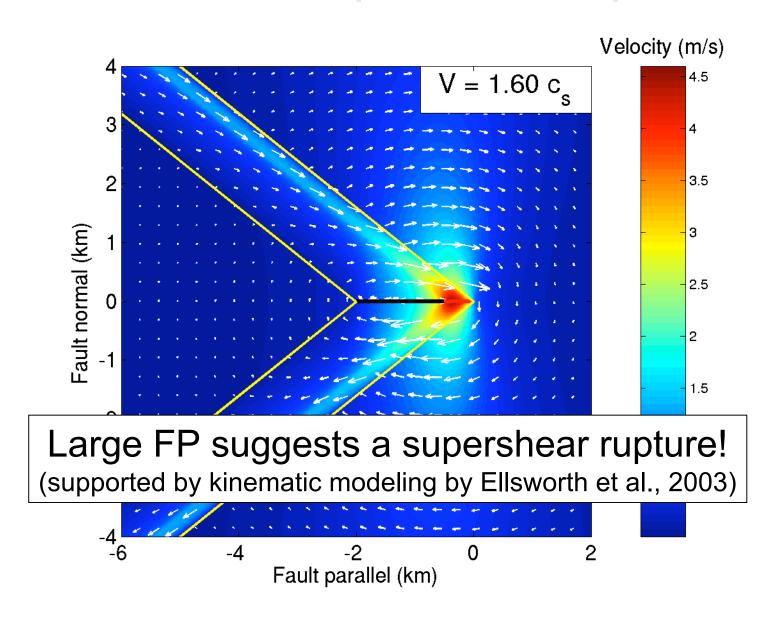
3km N of fault



#### The Puzzle

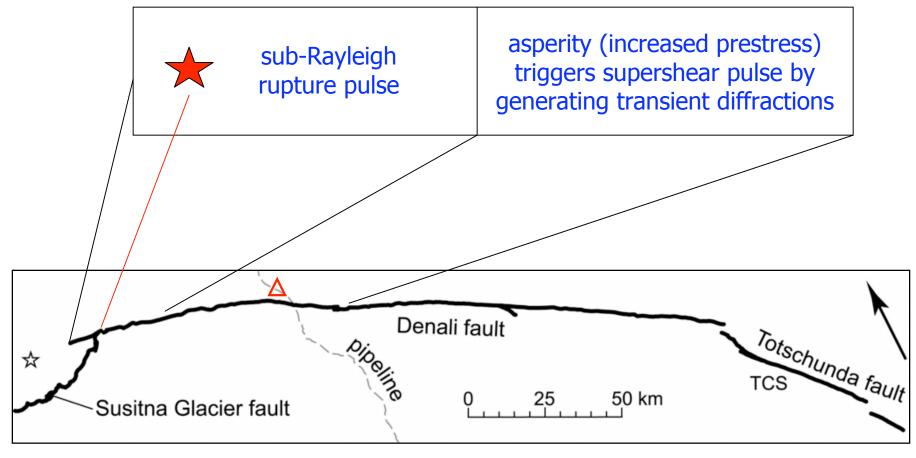


#### FP/FN and Supershear Ruptures



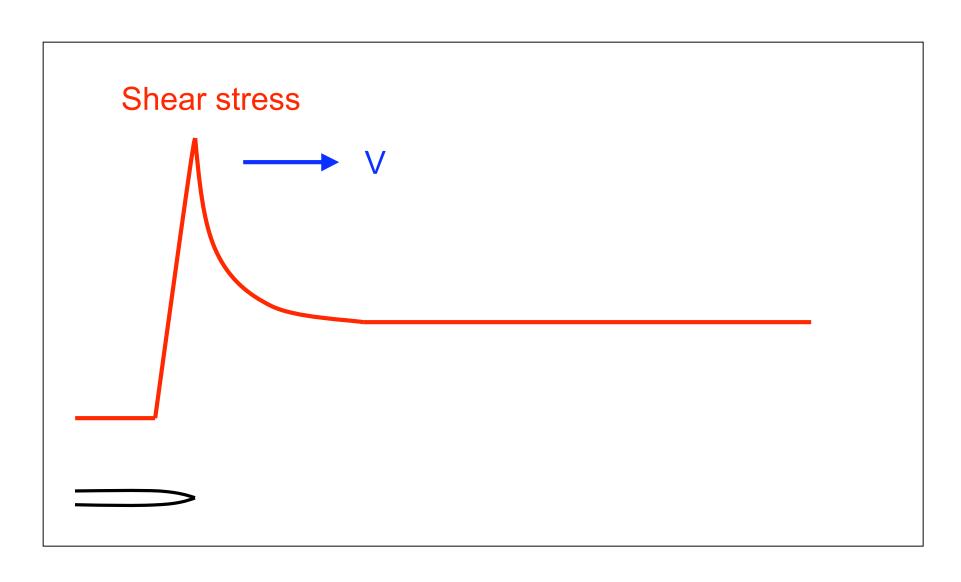
## A Spontaneous Dynamic Rupture Model

Extremely simple model – no depth dependence

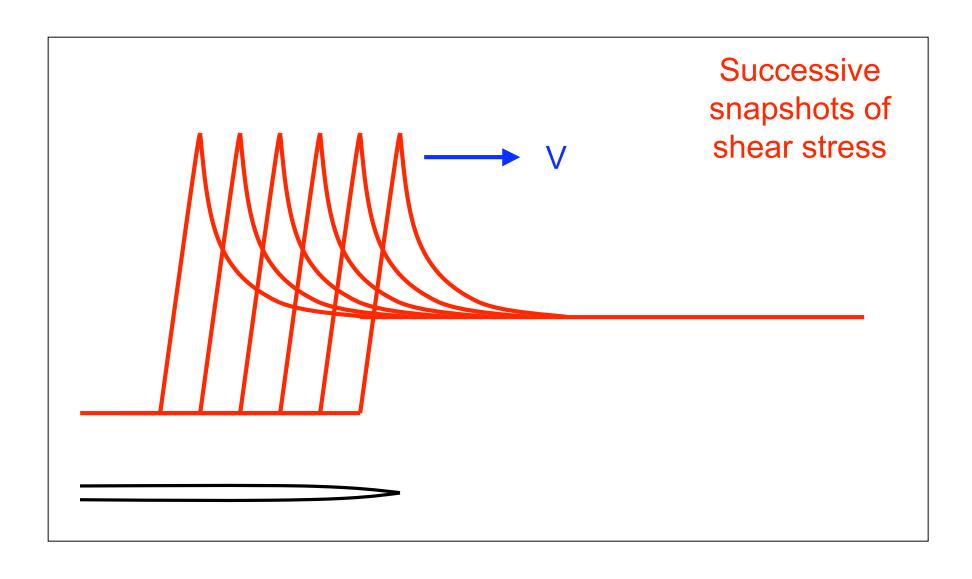


(Dunham and Archuleta, BSSA, 2005)

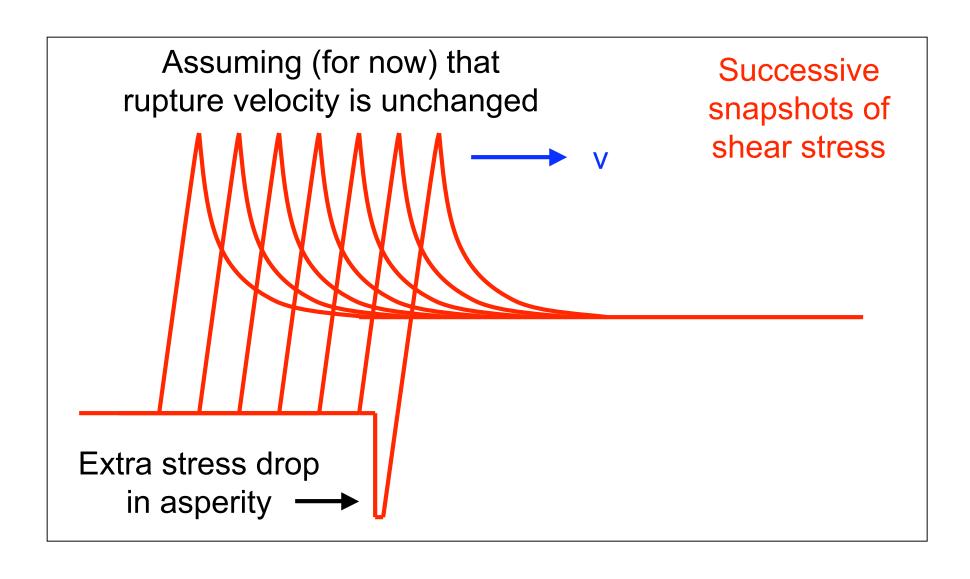
# Rupture of an Asperity: 1. Steady state rupture



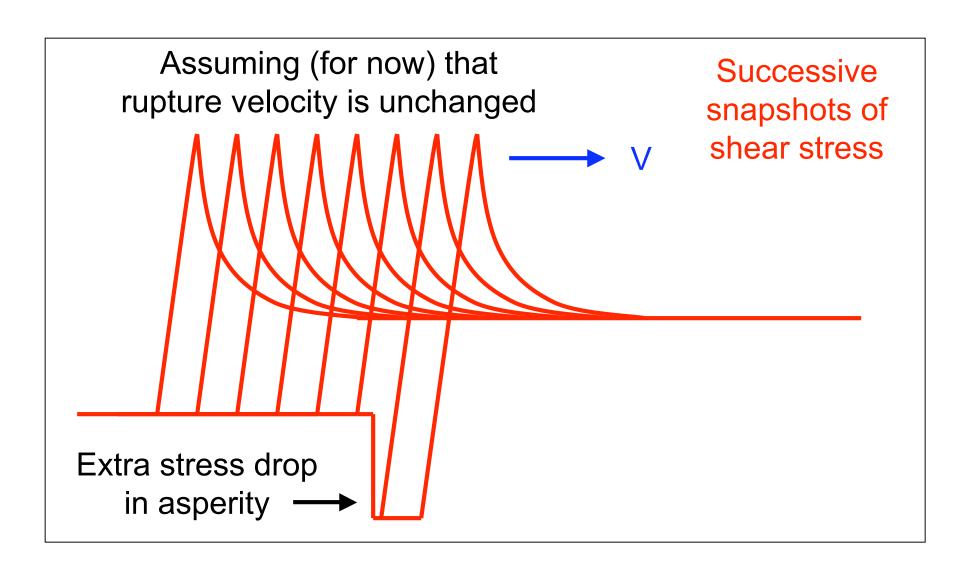
# Rupture of an Asperity: 1. Steady state rupture



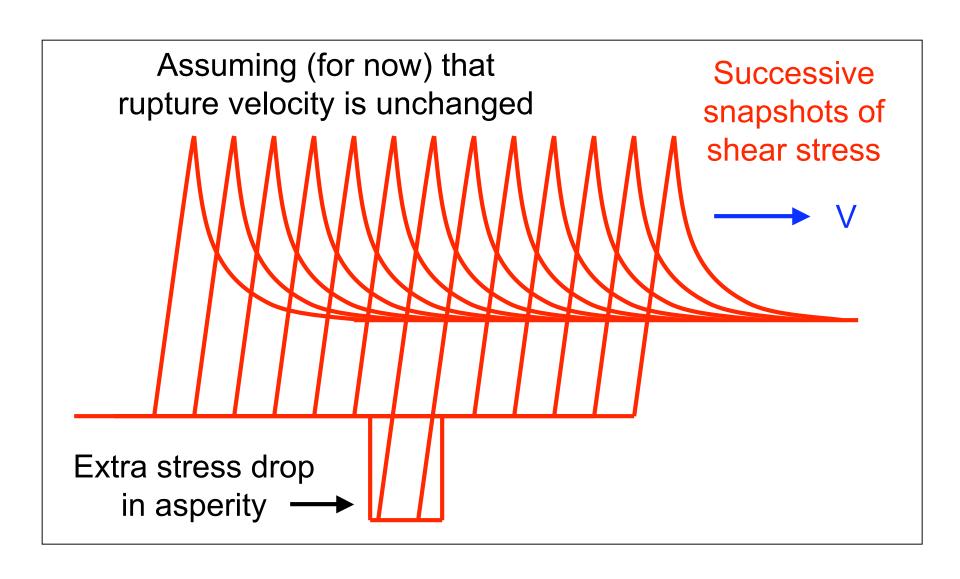
# Rupture of an Asperity: 2. Extra stress drop in asperity



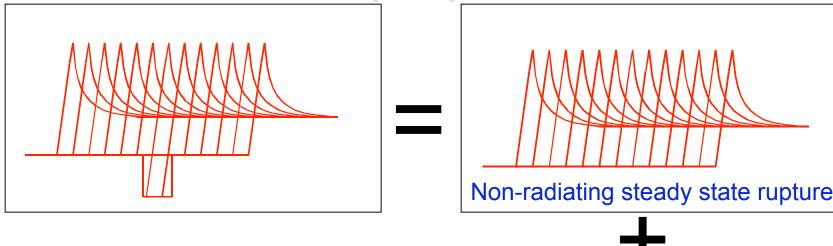
# Rupture of an Asperity: 2. Extra stress drop in asperity



# Rupture of an Asperity: 2. Extra stress drop in asperity



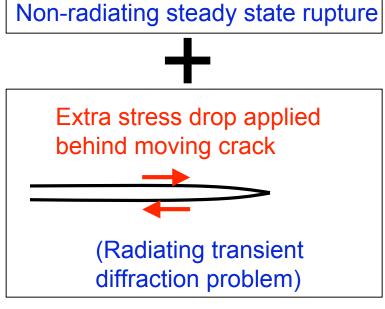
## Rupture of an Asperity: 3. Superposition



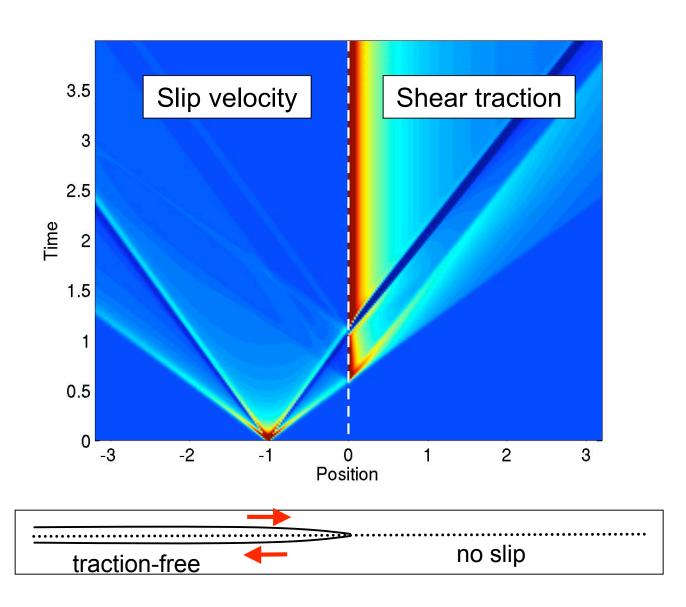
Note: Superposition only works when assuming constant rupture speed.

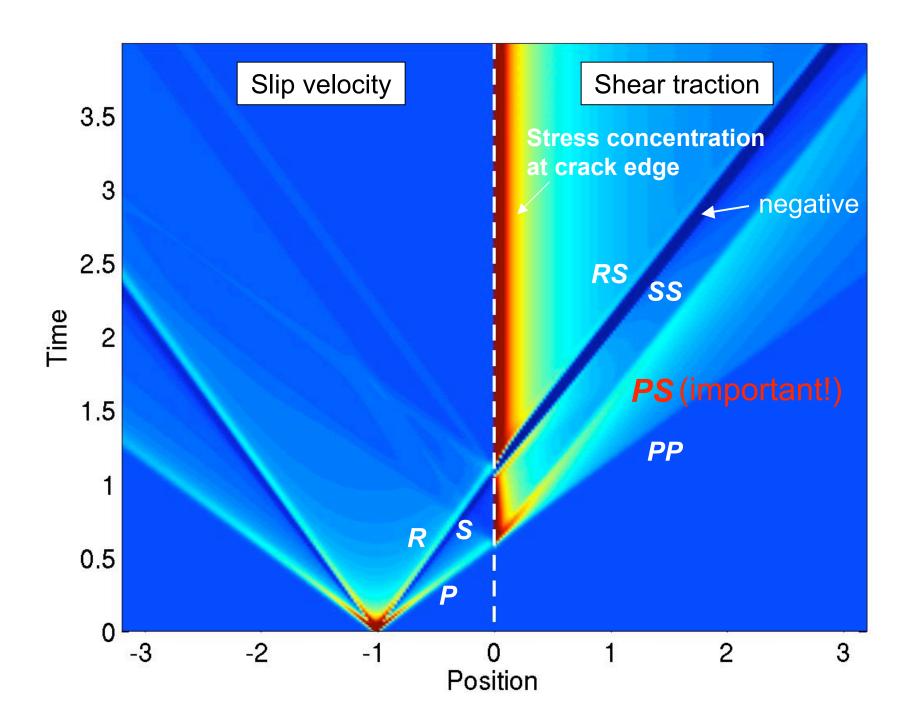
Relax this condition by determining the dynamics with a friction law

Transient waves will drive rupture at faster (even supershear) speeds

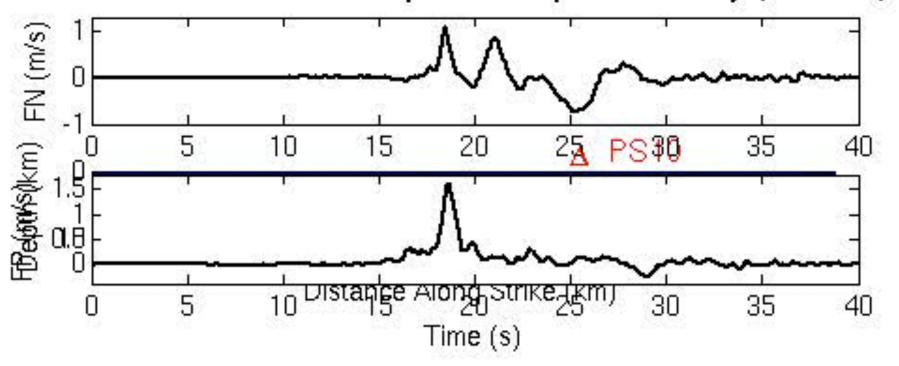


## Transient Diffractions: Step-function Stress Drop Behind Stationary Crack

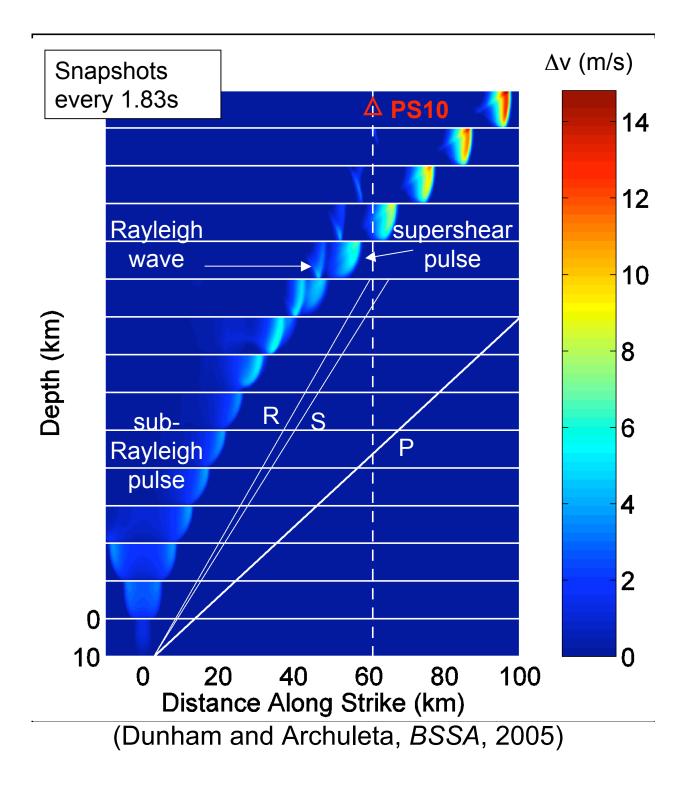


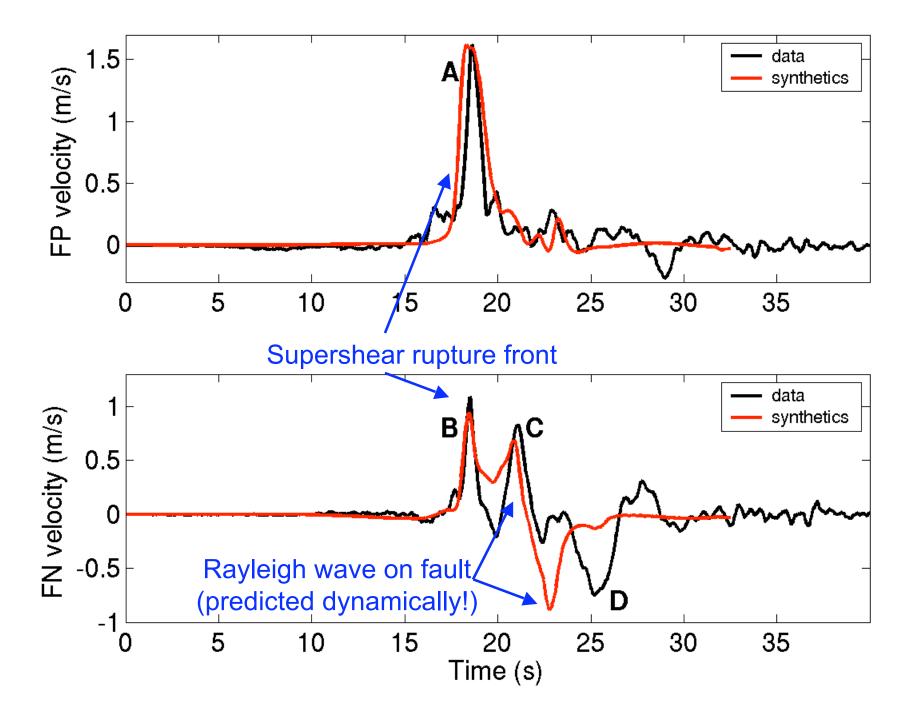


#### 2002 Denali Fault Earthquake - Rupture History (Model II)









#### Conclusions

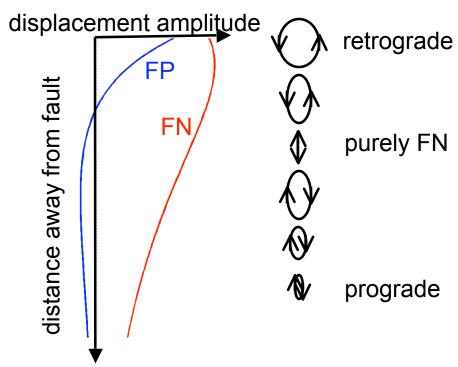
- Supershear ruptures do occur in nature
  - Denali Fault: 30km before PS10 to 30km after PS10
  - Have we really seen everything in only 50 years?
- Supershear records contain exact record of fault slip history
- Supershear generates large velocities far off of fault
  - Cause of extensive damage in Turkey?
  - Bay Bridge engineers already concerned about FP
  - Do our current design standards take this into account?



## Why is the Rayleigh wave absent from FP?

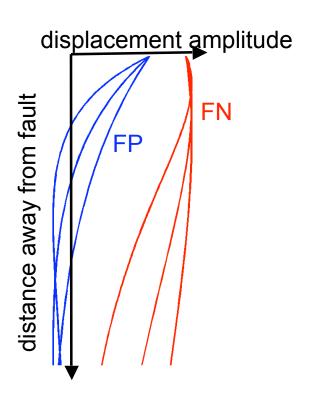
What we observe is a *pulse* (i.e., a superposition of harmonic Rayleigh waves of various wavelengths)

Examine harmonic Rayleigh wave (length scale is wavelength):



## Why is the Rayleigh wave absent from FP?

Superposition of various wavelengths causes destructive interference of FP (but not FN) component!



(pointed out by Michel Campillo)

#### Characterizing the Excited Waves

Decompose displacement field

$$u = u^p + u^s$$

$$\left(-\frac{1}{c_p^2}\frac{\partial^2}{\partial t^2} + \nabla^2\right)u^p(t, x, y) = 0 \qquad \nabla \times u^p = 0$$
Dilatational

$$\left(-\frac{1}{c_s^2}\frac{\partial^2}{\partial t^2} + \nabla^2\right)u^s(t, x, y) = 0 \qquad \nabla \cdot u^s = 0$$
Shear

#### **Excitations as Plane Waves**

Solutions of governing (wave) equations are just plane waves

$$e^{i(k_x x + k_y y - \omega t)}, \omega = kc$$

Let's parameterize them by

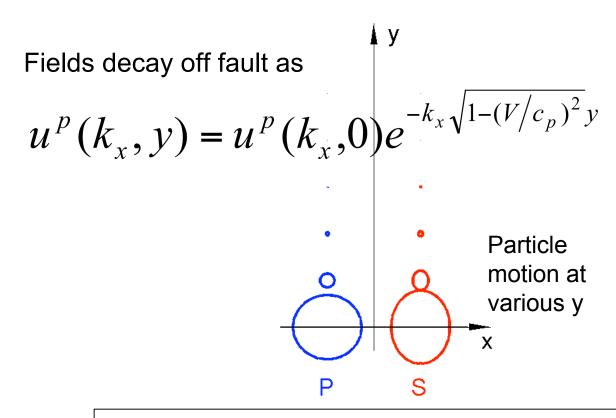
1. Along-fault phase velocity 
$$\omega/$$

1. Along-fault phase velocity 
$$\omega/k_x$$
 2. Along-fault wavelength  $2\pi/k_x$ 

Steady-state motion along fault requires

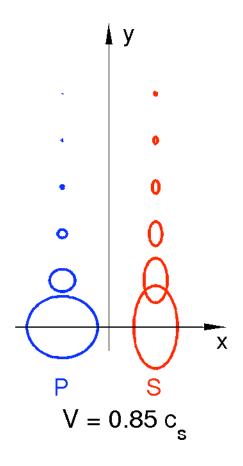
$$\omega/k_{x} = V$$
 (steady-state velocity)

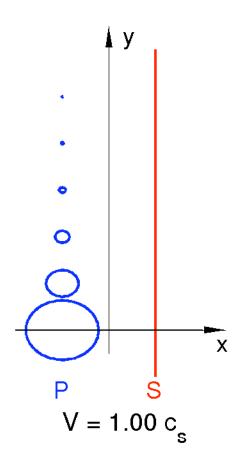
#### Inhomogeneous Waves (V<c)



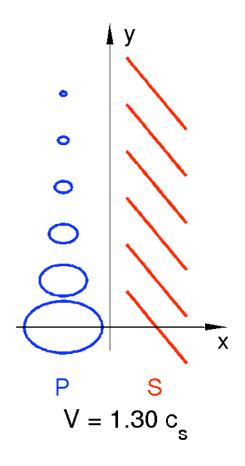
- 1. High frequency information is quickly lost as observation point moves away from fault
- 2. No far field radiation

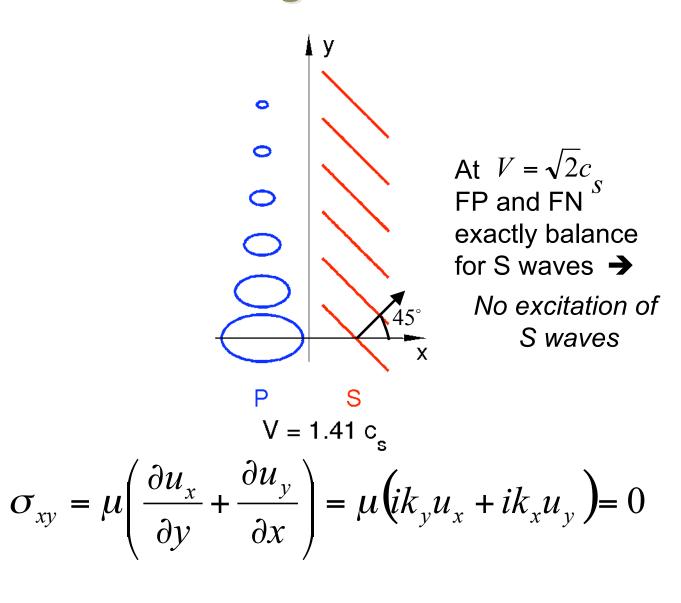
### Inhomogeneous Waves (V<c)

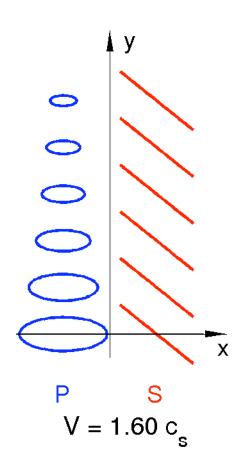




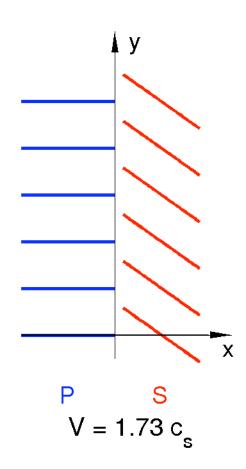
S waves start to radiate (amplitude independent of y)

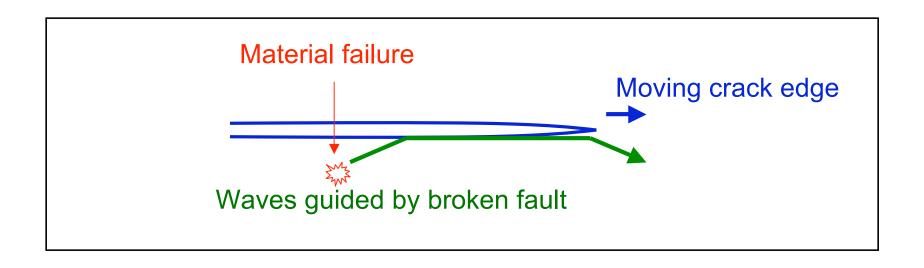


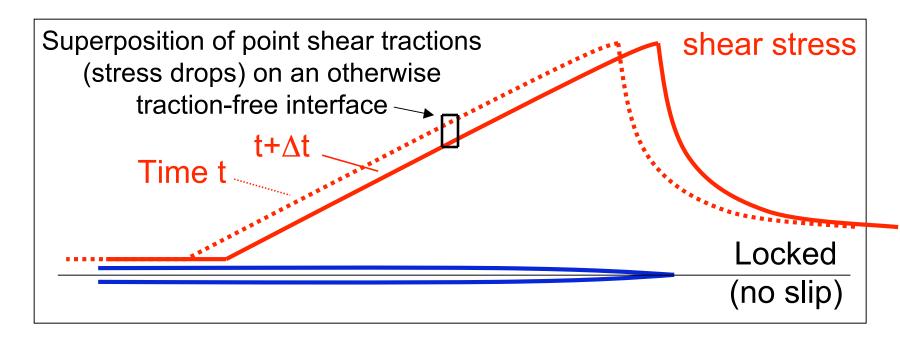




#### Grazing P wave + Radiating S wave

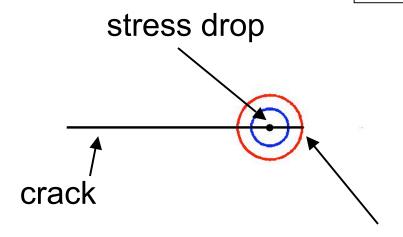






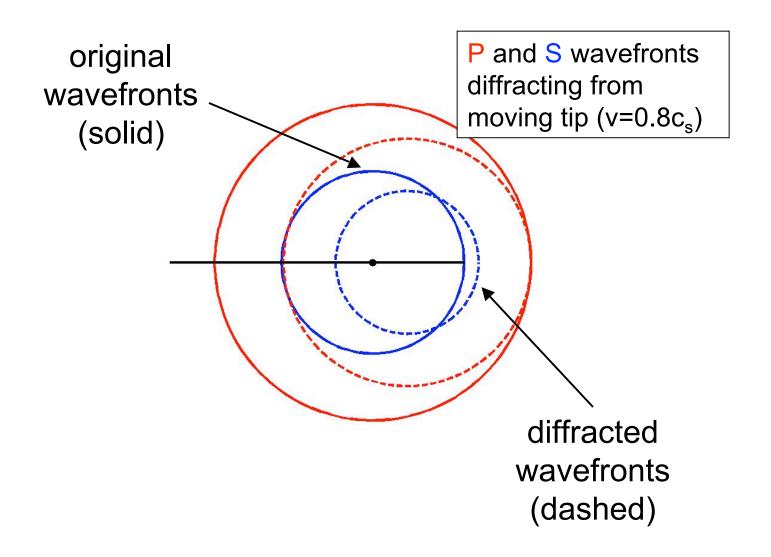
### Crack Edge Diffraction

P and S wavefronts diffracting from moving tip (v=0.8c<sub>s</sub>)

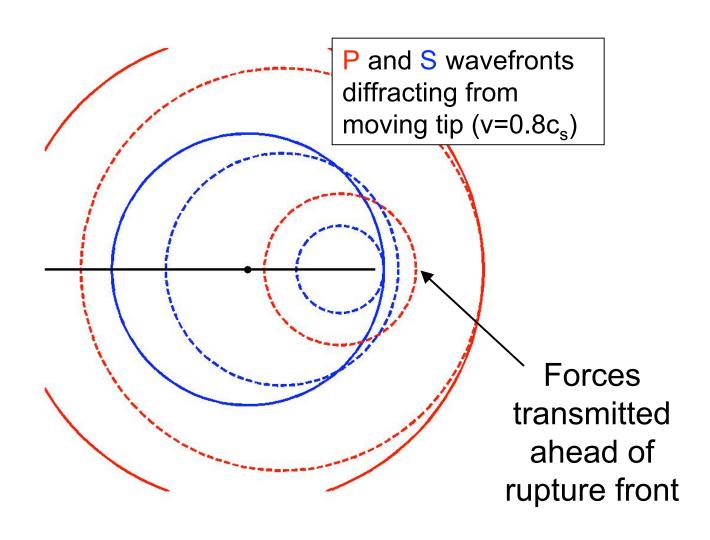


P-wave overtakes crack tip

### Crack Edge Diffraction



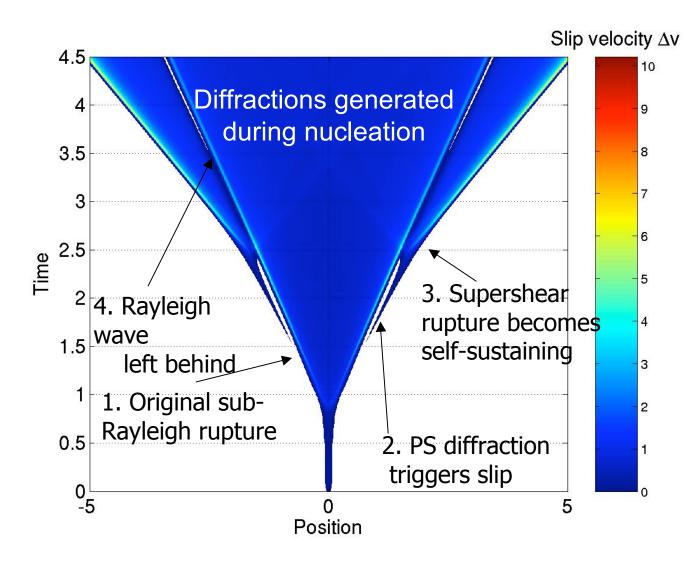
## Crack Edge Diffraction



#### Nucleation

During nucleation, crack tips are close together and waves bounce around between them

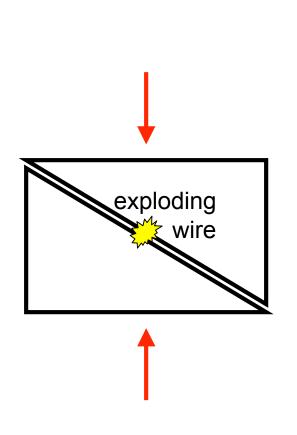
#### Multiple Diffractions Between Crack Tips

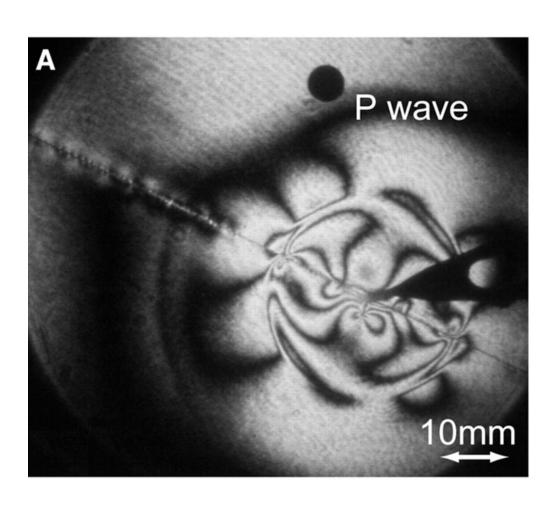


Predicted theoretically by Burridge (1973) and observed numerically by Andrews (1976)

## **Experimental Observation**

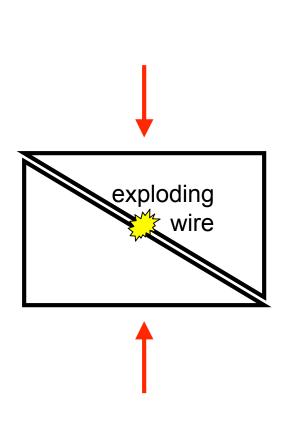
Laboratory experiments of Xia and Rosakis (2004)

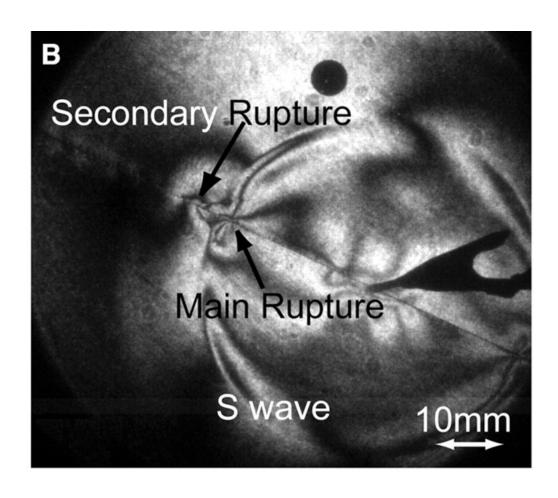




### **Experimental Observation**

Laboratory experiments of Xia and Rosakis (2004)





### **Experimental Observation**

Laboratory experiments of Xia and Rosakis (2004)

